ENGINEERING GEOLOGY

Syllabus of Engineering Geology INS1122, Spring Semester, 2014

Lecturer: Assistant Prof. Dr. Ömer Ündül – İstanbul University (İ.Ü.) Geological Engineering Dept. **Lecture:** Monday, 13:00 – 15:50

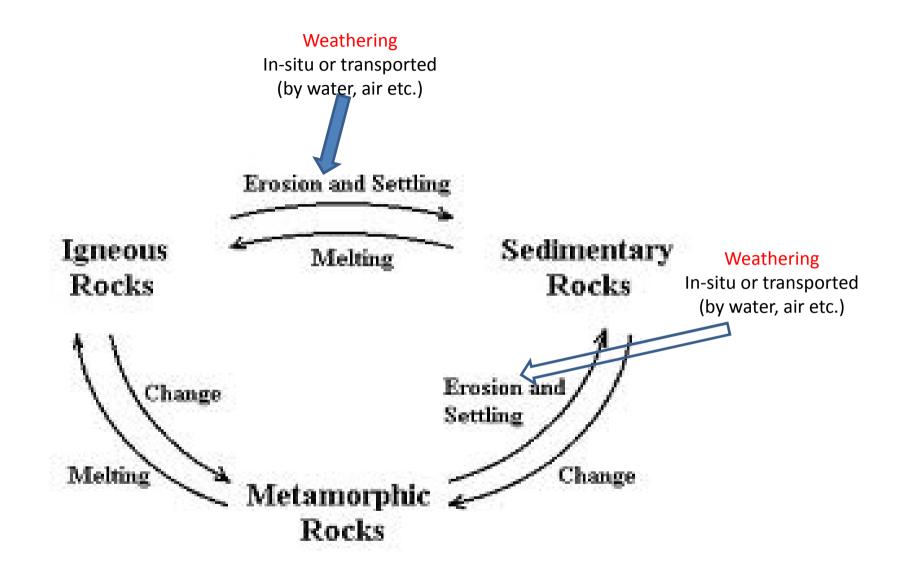
e-mail: oundul@istanbul.edu.tr,

http://aves.istanbul.edu.tr/SorguDokumanlarim.aspx?Sorgu=3217

Textbook: Geology for Civil Engineers by Marcus Matthews, Noel Simons and Bruce Menzies. A publ. of Thomson Telford.

Engineering Geology, Bell F.G, Second edition 2007, Butterworth-Heinemann, Elsevier

- 1) Introduction, the role of geology and engineering geology in civil engineering
- 2) The structure of the Earth's crust
- 3) Minerals and Rocks, classifications
- 4) Engineering properties of rocks
- 5) Topographic maps, geological maps and cross sections
- 6) Engineering geological maps
- 7) Earthquakes
- 8) Mass movements
- 9) Stability of rock slopes
- 10) Engineering geology in dams
- 11) Tunnelling
- 12) Groundwater
- 13) Natural and dimension stones
- 14) Environmental geology



Simplified rock cycle

WEATHERING

Weathering is the process of alteration and breakdown of rock and soil materials at and near the Earth's surface by physical, chemical and biological processes to form clay, iron oxides, and other weathering products.

Weathering is the process of physical disintegration and chemical decomposition of the rock

MECHANICAL / PHYSICAL

Disintegration of rocks and minerals by a physical or mechanical process

CHEMICAL

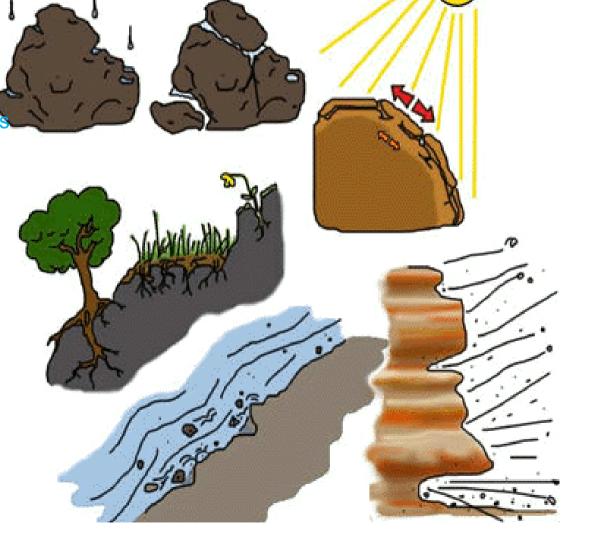
Chemical alteration or decomposition of rocks and minerals

BIOLOGICAL

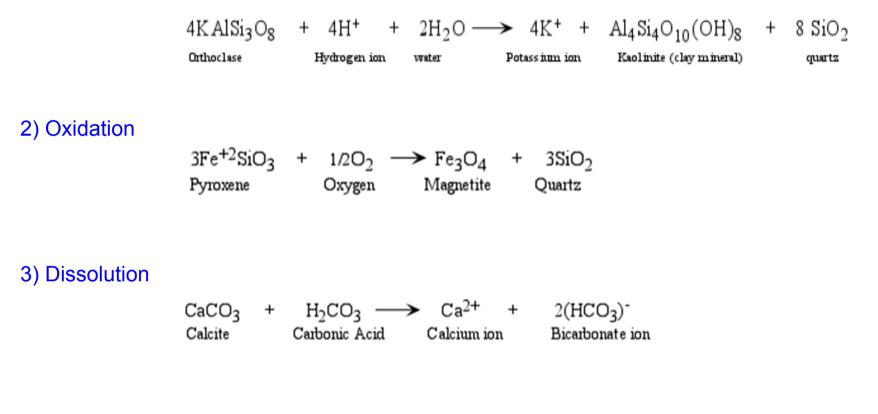
Roots, animal activity etc

WEATHERING - MECHANICAL / PHYSICAL Mechanical weathering

- 1) Development of Joints: Stress relief_exfoliation*, tectonics
- 2) Crystal Growth
- 3) Frost Action:
 - Freeze-thaw, frost wedging etc.
- 4) Changes in temperature:
 - Wet-dry cycles, slope aspect
- 5) Plant roots
- 6) Abrasion
- 7) Exfoliation **



1) Hydrolisis



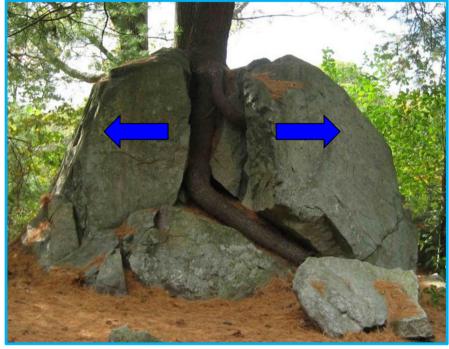
4) Hydration - Dehydration

2FeO∙OH →	Fe ₂ O ₃	+	H_2O
Goethite	Hematite		water

WEATHERING - BIOLOGICAL

1) Plant Roots

2) Organic Activities: algea, lichens, bacteria, movement of animals within rock (especially after completely weathered stage)



Modified from http://toms-randomthoughts.blogspot.com/2010/12/rocks.html



www.shutterstock.com

WEATHERING - LANDSCAPES



1) Aydın - Turkey



3) Antequera - Spain_The World Geography



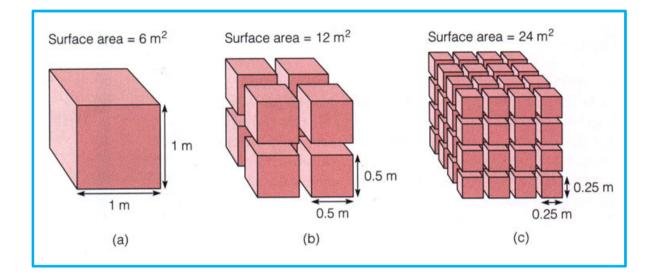
2) Kapadokya - Turkey



4) Grand Canyon - USA_National Geographic

WEATHERING - INFLUENCING FACTORS

1) Rock Type:



2) Particle Size:

3) Climate:

4) Time & Exposure

5) Topography:

6) Vegetation & Animal activities

WHY ARE WE INTERESTED IN WEATHERING?

1) Soil Development

2) Slope Stability

3) Tunnelling

4) Geological Materials

5) Agriculture etc.

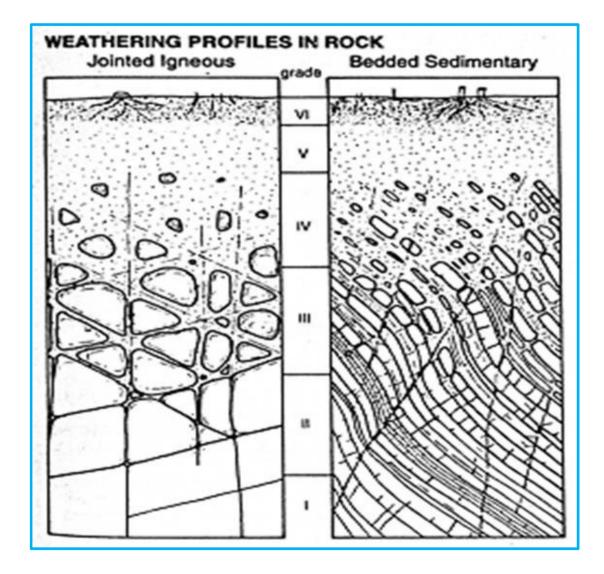
WEATHERING CLASSIFICATION

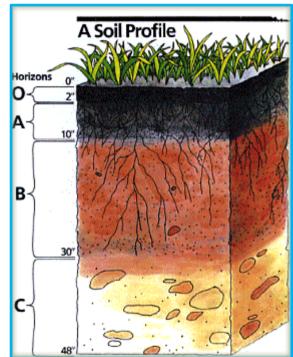
Weathering

grades

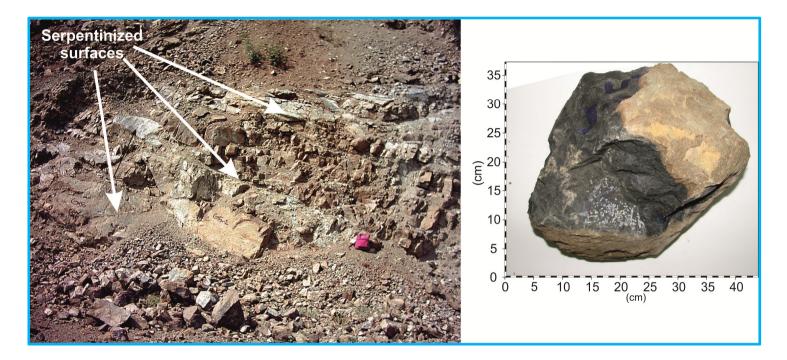
VI	Residual soil Soil derived by in situ weathering but having lost original texture and fabric
v	Completely weathered Considerably weakened . Slakes in water. Original texture apparent
IV	Highly weathered Large pieces can be broken by hand. Does not readily disaggregate (slake) when dry sample immersed in water.
ш	Moderately weathered Considerably weakened, penetrative discoloration. Large pieces cannot be broken by hand
П	Slightly weathered Slight discoloration slight weakening
I	Fresh Unchanged from original state

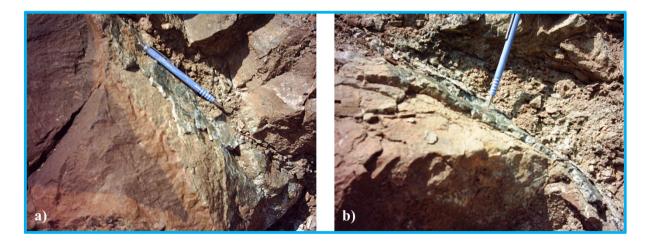
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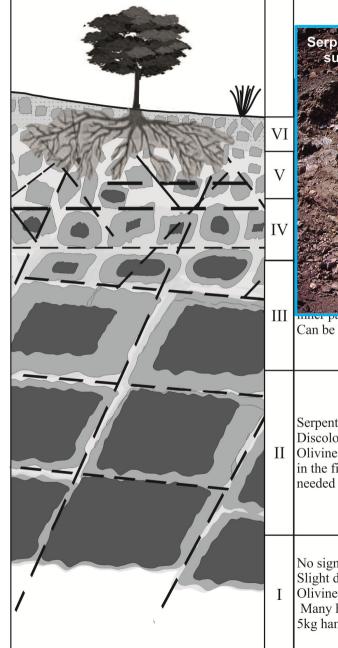


WEATHERING of ULTRAMAFIC ROCKS

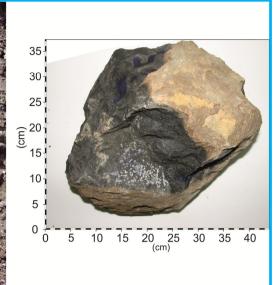




WEATHERING PROFILE of ULTRAMAFIC ROCKS







Can be broken by some hammer blows

Serpentinization along discountinuities. Discolouration due to iron oxide. Olivine and pyroxene minerals can be observed in the field. Many blows with a 5 kg hammer are needed to break the rock.

No sign of weathering. Slight discolouration along cracks. Olivines can be identified with naked eyes. Many hard blows are needed with a 5kg hammer to break a piece.

WEATHERING of ULTRAMAFIC ROCKS



WEATHERING of ULTRAMAFIC ROCKS

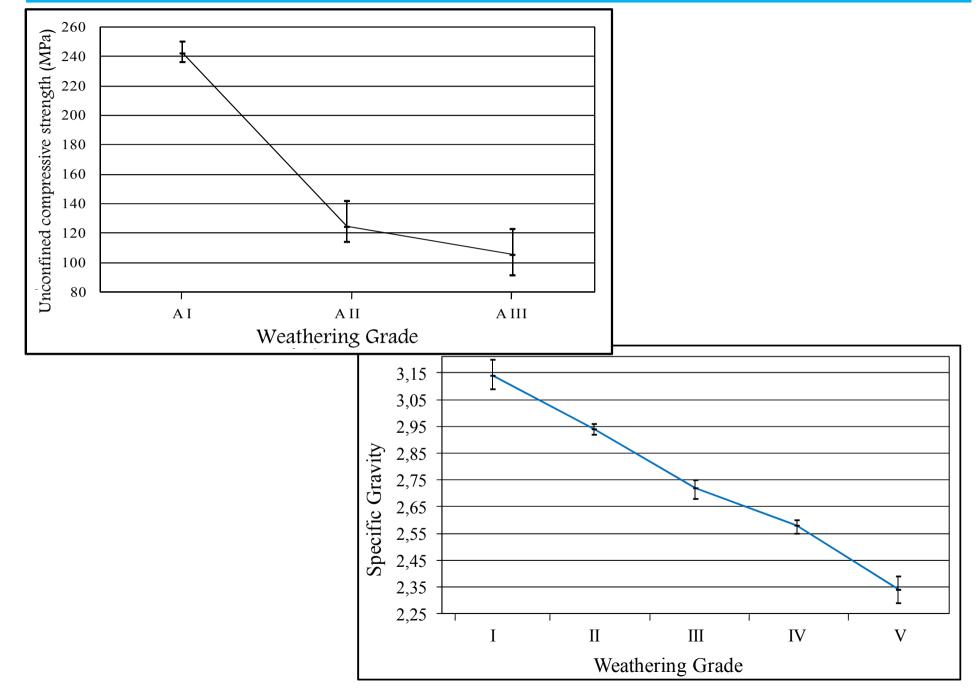


WEATHERING



WEATHERING of ULTRAMAFIC ROCKS





Grade	Density γ (kN m ⁻³)	Porosity n (%)	Unconfined compressive strength UCS (MPa)			
Dolerite ^a						
I—II	28.04	0.4	c1.60e ra n o 1.80 rding			
Ш	27.64	0.5	Change ratio according 83 — 160			
IV	26.96	1	58 - 83			
V	26.18	3.2	24 - 58			
Granodiorite ^b						
1	25.6-26.96	2.6-0.4	111 - 165			
11	25.7–26	3.3-5.9	60 - 97			
Ш	25.1-25.7	1.5–2.3	33 - 48			
IV	22.9-25.1	5.2-6.1	8 - 24			
V	19.8	24	0.1			
VI	14.7	44				

Due to weathering Colours of the rock changes Porosity increases Grains get weaker The bonds between grains weakens Increase of micro-cracks

Consequently,

Rock loses the strength Permeability may subject to change Shear strength values decreases along discountinuities Specific gravity decreases

<u>The depth of weathered zone</u> <u>and</u> weathering grades should be determined</u>

Rocks can be used as

- 1) Construction material
 - a) Fill materail (highways, dams etc.)
 - b) Industrial use as dimensional stones or aggregates

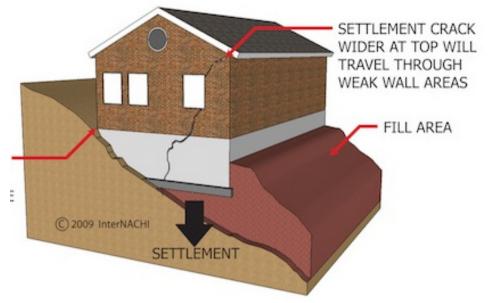
2) Foundation material

- a) Building, viaduct, dam etc.
- b) Tunnel, dam, slopes etc.

Engineering Geology for Civil Engineers

ENGINEERING PROPERTIES of ROCKS Construction material

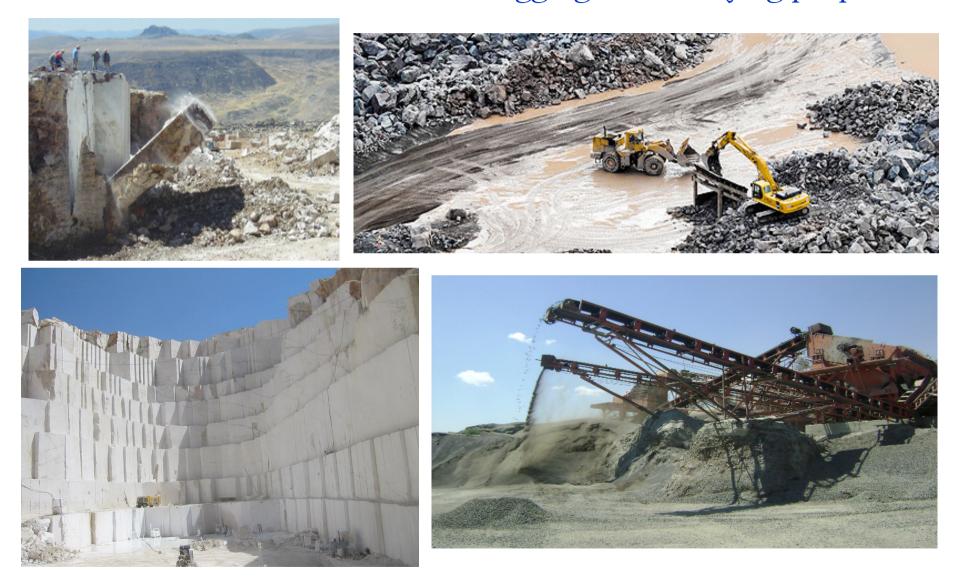


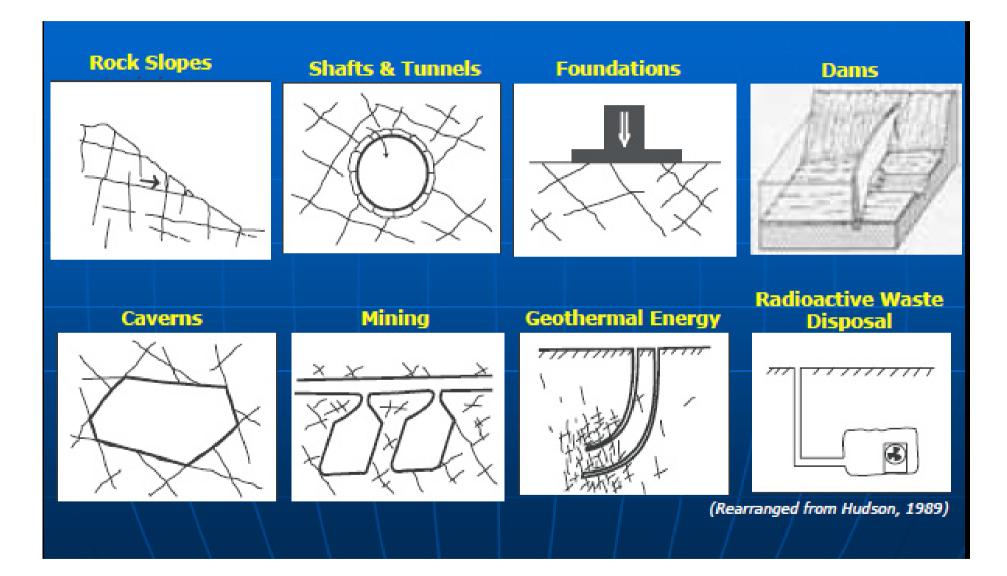






ENGINEERING PROPERTIES of ROCKS Natural stone, dimension stone, aggregate for varying purposes









- 1. Density, specific gravity, porosity, and void ratio
- 2. Rock strength
 - Definition of stress, normal and shear stress
 - Rock strength: tensile, shear, and compressive
 - Failure criteria: tensile, compressive
 - Failure criterion: shear failure
- 3) Discontinuity properties of rocks

Rock are significant for two major reasons in engineering

- 1) Building materials for constructions
- 2) Foundations on/in which the constructions are build

The consideration of rocks as construction material the engineers concern about

- a) Density (to some extent)
- b) Strength
- c) Durability (the ability of a material to resist degradation by mechanical or chemical agents)

For the consideration of rocks as the construction foundation the engineers concern about

- a) Density
- b) Strength
- c) Deformation
- d) Discontinuity properties

1. Density, specific gravity, porosity, and void ratio

Specific gravity

 $=\frac{\rho_s}{\rho_w}$ G_{s}







Specific Gravity is determined in crushed (powdered) specimen and specific gravities of most rock-forming minerals range from 2.5 to 2.8, although some heavier minerals are from 3 to 5 and/or higher

Density, specific gravity, porosity, and void ratio

Density

 $\rho = \frac{m}{v}$

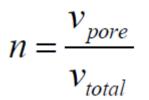
Density is obtained from representative piece of the rock itself.



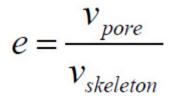


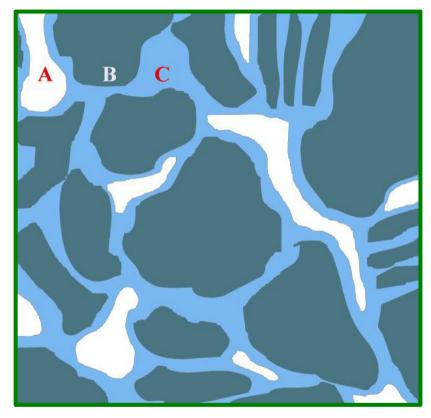
Density, specific gravity, porosity, and void ratio

Porosity



Void ratio



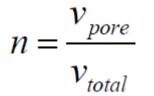




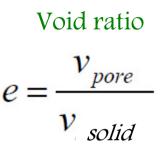
Schematic representation of porous medium indicating relationship between air (A), solid (B) and water (C).

Density, specific gravity, porosity, and void ratio

Porosity







Porosity depends largely on rock origin. Slowly cooling igneous magma results in relatively nonporous rock, whereas rapid cooling associated with escaping gases yields a porous mass.

Sedimentary rocks depend on amount of cementing materials present and on size, grading, and packing of particles.

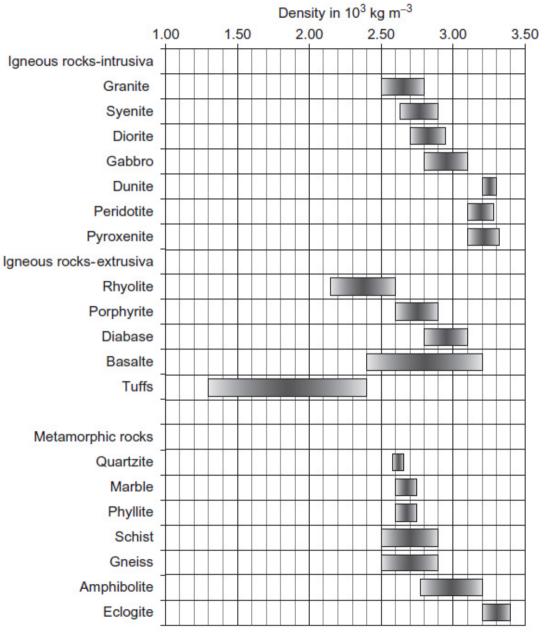


FIGURE 4.3 Mean range of density for igneous and metamorphic rocks.

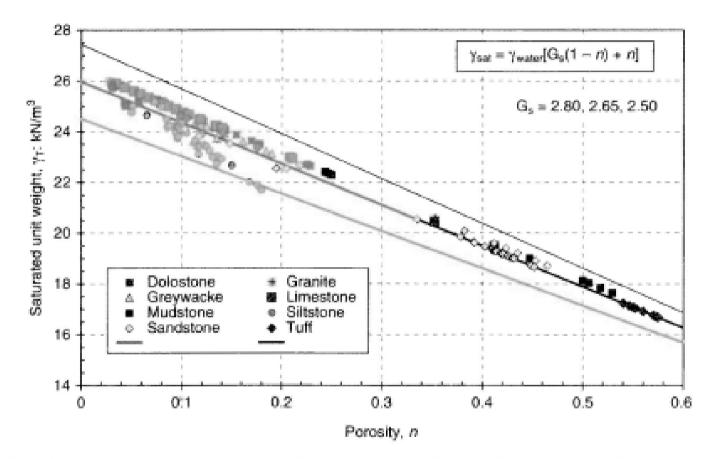
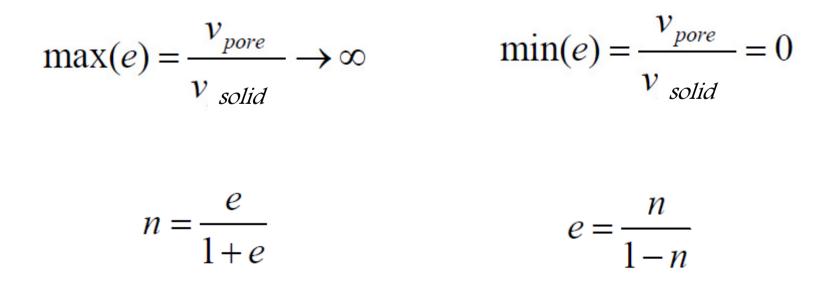


Fig. 9.4 Saturated rock unit weight in terms of porosity and specific gravity

The limits of the value of porosity

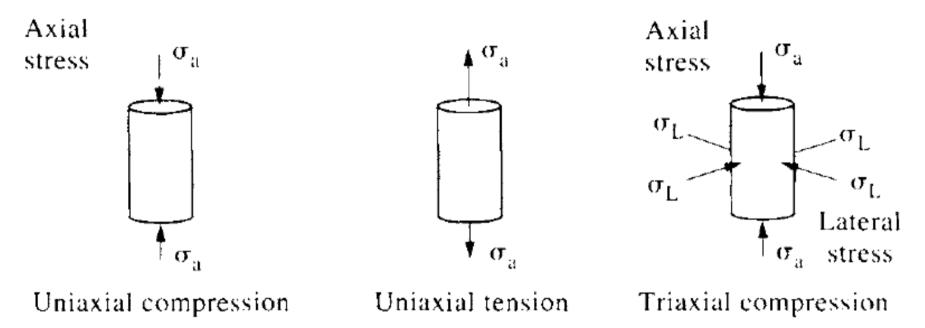
$$\max(n) = \frac{v_{pore}}{v_{total}} = 1 \qquad \min(n) = \frac{v_{pore}}{v_{total}} = 0$$

The limits of the value of void ratio

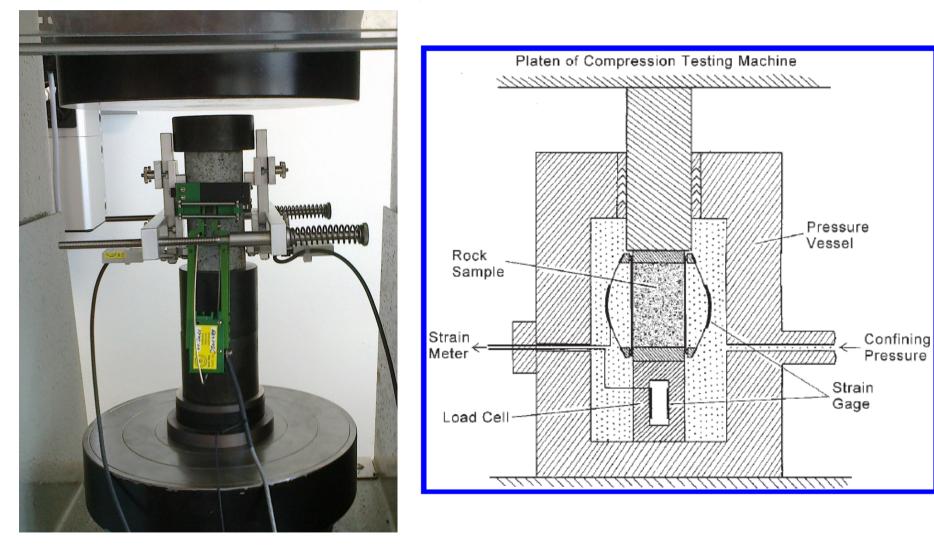


2. Rock strength,

is the peak stress to break down the rock sample *in a uniaxial or triaxial* compression process (modified from Farmer, 1982)

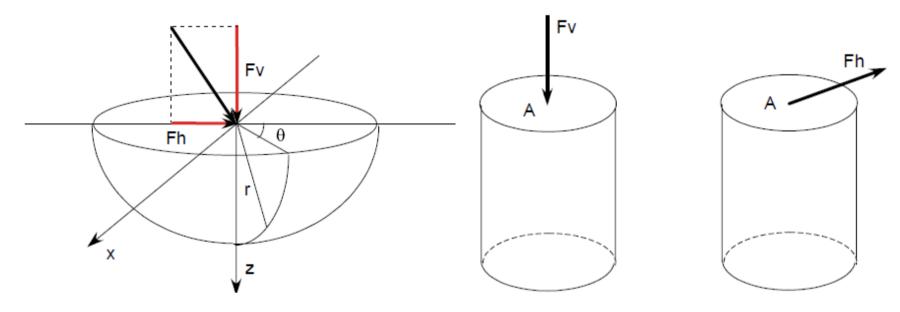


Rock strength, *is the peak stress* to break down the rock sample *in a uniaxial or triaxial* compression process (modified from Farmer, 1982)



What is stress?

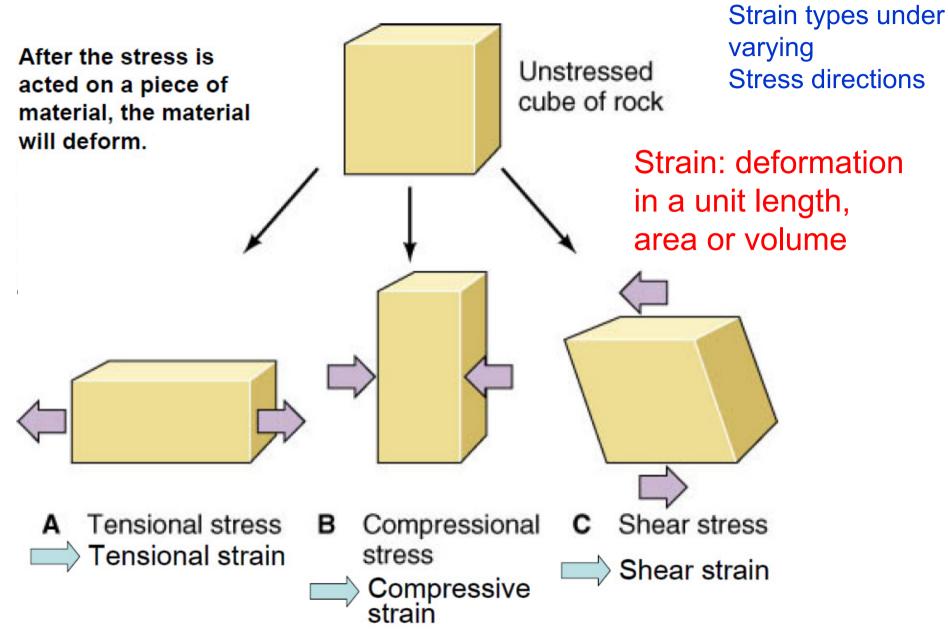
Stress is the force per area applied on the object



 $\sigma_n = \frac{F_v}{4} \qquad \sigma_s = \sigma_n$ = normal stress σ_s = shear stress

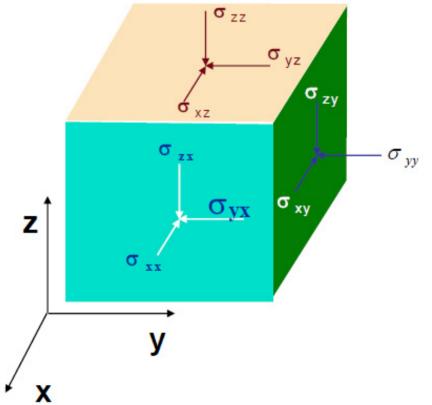
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ENGINEERING PROPERTIES of ROCKS



Actually, since the force is a vector, and can act on any direction, it can be dissolved into 3 orthogonal directions (e.g., x, y, z, in Cartesian coordinate); a surface in or on the solid can face any direction, too. The normal of that surface is also a function of x, y, z. Consequently, a complete stress should be a tensor with 9 elements.

$$\mathbf{\sigma} = \begin{bmatrix} \sigma_{xx} & \sigma_{xy} & \sigma_{xz} \\ \sigma_{yx} & \sigma_{yy} & \sigma_{yz} \\ \sigma_{zx} & \sigma_{zy} & \sigma_{zz} \end{bmatrix}$$



Rock strength

As the stress can be divided into normal stress (compressive and tensile) and shear stress. Similiarly the rock should also have;

compressive strength (CS)

shear strength (τ_0)

tensile strength (T)

For the same kind of rock usually

 $CS > \tau_0 > T$

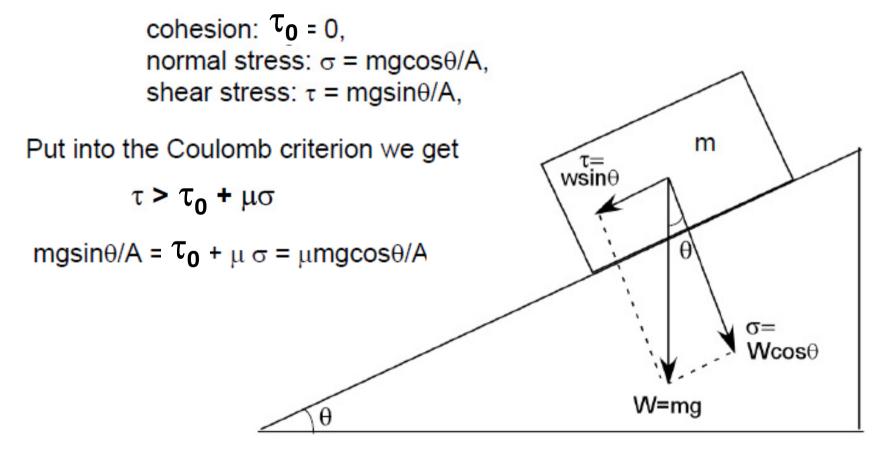
Rock strength and failure criteria

However, the shear failure is not as simple as the tensile and compressive failure criteria. (<u>Because it involves both</u> <u>normal and shear stresses</u>). Laboratory experiments in rock mechanics shows that the shear failure is more complicated and follows the Coulomb Criterion stated as

 $\tau > \tau_0 + \mu \sigma$

The Coulomb Criterion shows that shear stress tends to cause failure across a plane, it is resisted by the sum of the shear strength (cohesion) of the material and a constant times the normal stress across the plane. This constant is the <u>coefficient of friction</u>. When the shear stress is equal to or greater than that sum, shear failure occurs.

Let's illustrate the Coulomb criterion for shear failure by a very common test (see the sketch below). First, assume that there is no inherent shear strength (cohesion) between the block (with a mass of m) and the underlying slope. The contact area is A, and the slope on which the block is possible sliding has a dip angle of θ . We have the following relations:



Rock strength and failure criteria

Tensile failure occur

|σ**| > T**

Compressive failure occur |σ| > CS

Shear failure occur (Coulomb criterion) $\tau > \tau_0 + \mu \sigma$

Strain

Change in shape or size of an object in response to an applied stress.

= Deformation

Three Types of Strain

•Elastic

•Ductile (Plastic)

•Brittle (Rupture)

Elastic Deformation

A temporary change in shape or size that is recovered when the applied stress is removed.

If the response of the material to the load/unload is instantaneous, it is a <u>pure</u> <u>elastic</u> material;

If the response of the material to the load/unload needs finite time, it is a visco-elastic material;

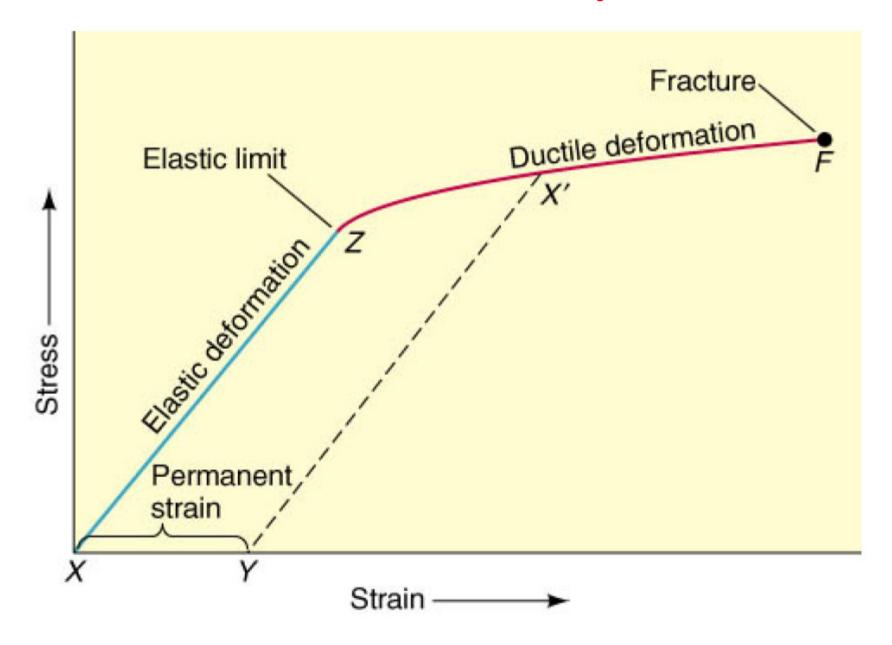
Ductile (Plastic) Deformation

 A permanent change in shape or size that is not recovered when the stress is removed.

• i.e. it flows or bends

BRITTLE DEFORMATION

The failure of a specimen without any visible plastic deformation



Young's modulus E

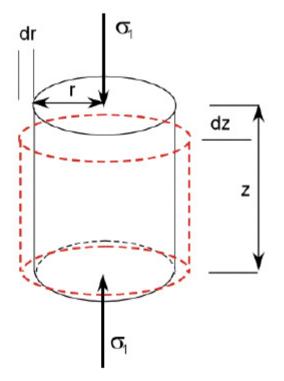
Young's modulus is the stress needed to compress the solid to shorten in a unit strain.

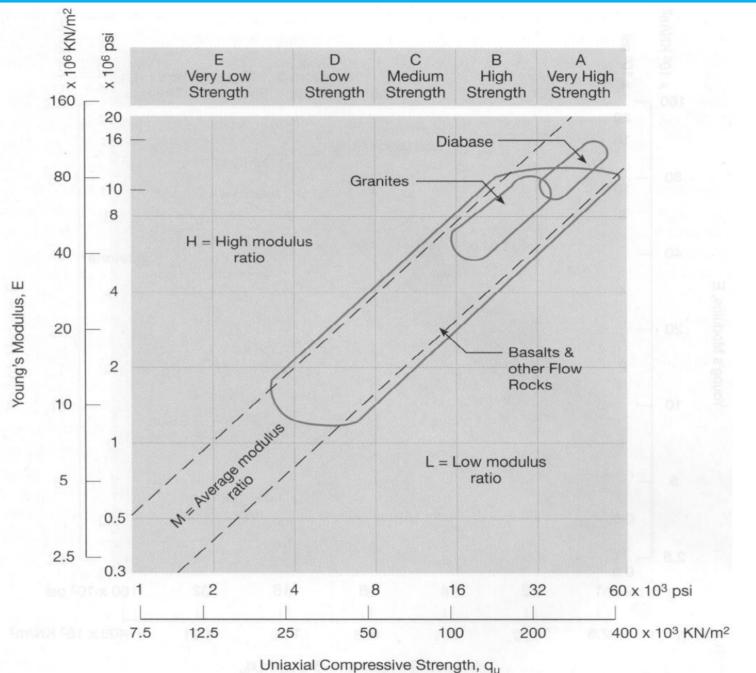
$$E = \frac{\sigma_1}{\Delta z \, / \, z}$$

Poisson's ratio v

Poisson's measures the relativity of the expansion in the lateral directions and compression in the direction in which the uni-axial compression applies.

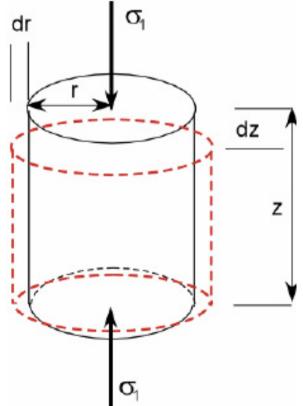
$$\nu = -\frac{\Delta r \,/\, r}{\Delta z \,/\, z}$$



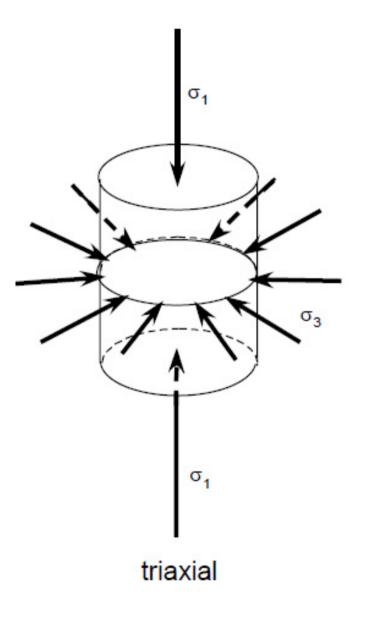


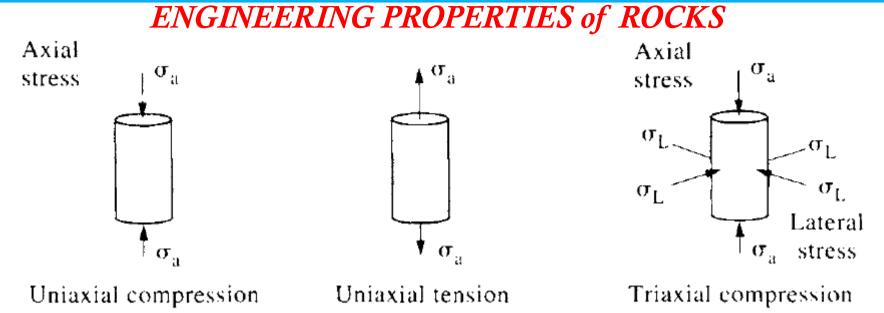
ENGINEERING PROPERTIES of ROCKS Uni-axial Compression Test

In a uni-axial compression test, the direction of the load is called the maximum principal direction and there are no other loads (forces) working on other direction.



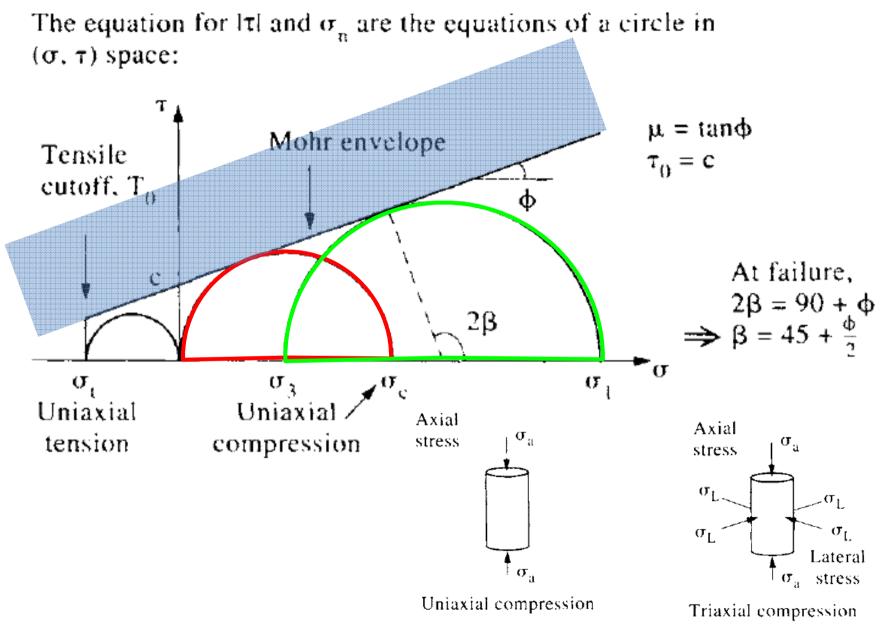
In a triaxial compression test, the direction of the load is called the maximum principal direction and the direction of the confining pressure applied is the minimum principal direction.





Rock fails at a critical combination of normal and shear stresses:

 $\tau_{0} = \text{cohesion } \mu = \text{coeff. of friction}$ $\tau_{0} = \frac{1}{2}(\sigma_{1} - \sigma_{3}) \sin 2\beta$ $\sigma_{n} = \frac{1}{2}(\sigma_{1} + \sigma_{3}) + \frac{1}{2}(\sigma_{1} - \sigma_{3}) \cos 2\beta$



Rock category	General description	Rock	Unconfined compressive strength, $\sigma_c^{(1)}$ (MPa)
A	Carbonate rocks with well-	Dolostone	33-310
	developed crystal cleavage	Limestone	24 - 290
		Carbonatite	38 - 69
		Marble	38 - 241
		Tactite-Skarn	131 – 338
В	Lithified argillaceous rock	Argillite	29-145
		Claystone	1 - 8
		Marlstone	52 - 193
		Phyllite	24 - 241
		Siltstone	10 - 117
		Shale ⁽²⁾	7 – 35
		Slate	145 - 207
С	Arenaceous rocks with strong crystals and poor cleavage	Conglomerate	33 - 221
		Sandstone	67 – 172
		Quartzite	62 - 379
D	Fine-grained igneous crystalline rock	Andesite	97 – 179
		Diabase	21 - 572
Ε	Coarse-grained igneous and metamorphic crystalline rock	Amphibolite	117-276
		Gabbro	124 - 310
		Gneiss	24-310
		Granite	14 - 338
		Quartz diorite	10 – 97
		Quartz monozonite	131 – 159
		Schist	10 - 145
		Syenite	179 – 427

 Table 7.1
 Typical range of unconfined compressive strength of intact rocks (AASHTO, 1989).

⁽¹⁾Range of unconfined compressive strength reported by various investigators.

⁽²⁾Not including oil shale.

TABLE 5.20

Suggested Hardness Classification for Intact Rock^a

Class	Hardness	Field test	Strength ^b tsf
Ι	Extremely hard	Many blows with geologic hammer required to break intact specimen	>2000
П	Very hard to hard	Hand-held specimen breaks with	2000-700
_		hammer end of pick under more than one blow	700–250
III	Moderate	Cannot be scraped or peeled with knife, hand-held specimen can lie broken with single moderate blow with pick	250-100
IV	Soft	Can just be scraped or peeled with knife. Indentations 1–3 mm deep shown in specimen with moderate blow of pick	100–30
V	Very soft	Material crumbles under moderate blow with sharp end of pick and can be peeled with a knife, but is too hard to hand-trim for triaxial test specimen	30–10

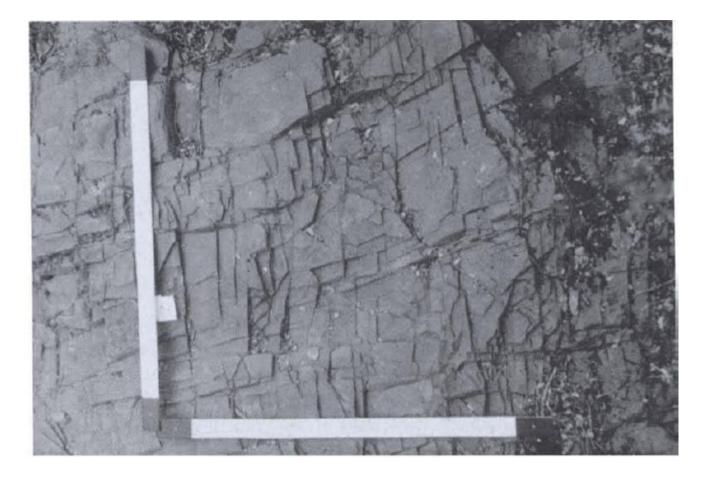
^a After ISRM Working Party, Int. Sec. Rock Mech., Lisbon, 1975.

^b Uniaxial compressive strength (Core Logging Comm., Bull. Assoc. Eng. Geol., XV, 295–328, 1978.

İstanbul University – Yıldız Technical University

Engineering Geology for Civil Engineers

ENGINEERING PROPERTIES of ROCKS 3) Discontinuity properties of rocks



ENGINEERING PROPERTIES of ROCKS 3) Discontinuity properties of rocks

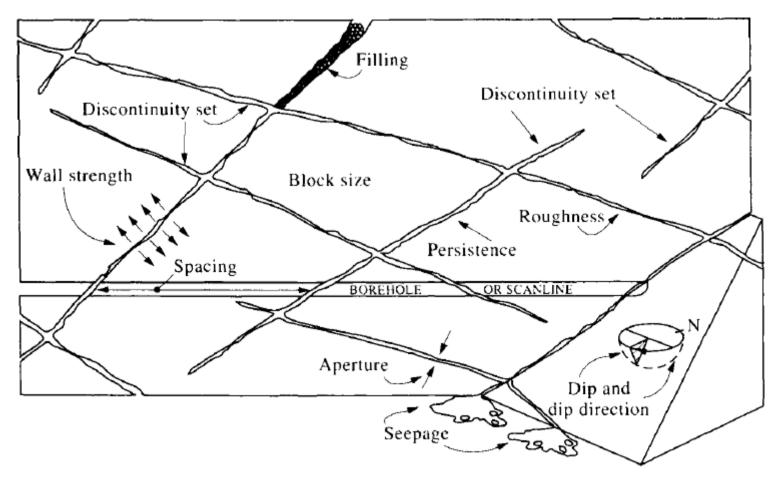


Figure 7.3 Schematic of the primary geometrical properties of discontinuities in rock (from Hudson, 1989).

Discontinuity properties of rocks

1. Orientation. The attitude of a discontinuity in space. It is described by the dip direction (azimuth) and dip of the line of steepest declination in the plane of the discontinuity.

2. **Spacing**. The perpendicular distance between adjacent discontinuities. It normally refers to the mean or modal spacing of a set of discontinuities.

3. Persistence: The discontinuity trace length as observed in an exposure. It may give a crude measure of the areal extent or penetration length of a discontinuity. Termination in solid rock or against other discontinuities reduces the persistence.

4. Roughness. The inherent surface roughness and waviness relative to the mean plane of a discontinuity. Both roughness and waviness contribute to the shear strength. Large scale waviness may also alter the dip locally.

5. Wall strength. The equivalent compressive strength of the adjacent rock walls of a discontinuity. It may be lower than rock block strength due to weathering or alteration of the walls. It is an important component of shear strength if rock walls are in contact.

ENGINEERING PROPERTIES of ROCKS Discontinuity properties of rocks

6. Aperture: The perpendicular distance between adjacent rock walls of a discontinuity, in which the intervening space is air or water filled.

7. Filling: The material that separates the adjacent rock walls of a discontinuity and that is usually weaker than the parent rock. Typical filling materials are sand, silt, clay, breccia, gouge, mylonite. It also includes thin mineral coatings and healed discontinuities such as quartz and calcite veins.

8. **Seepage**. The water flow and free moisture visible in individual discontinuities or in the rock mass as a whole.

9. Number of Sets: The number of discontinuity sets comprising the intersecting discontinuity system. The rock mass may be further divided by individual discontinuities. 10.Block Size: The rock block dimensions resulting from the mutual orientation of intersecting discontinuity sets, and resulting from the spacing of the individual sets. Individual discontinuities may further influence the block size and shape.

Discontinuity properties of rocks

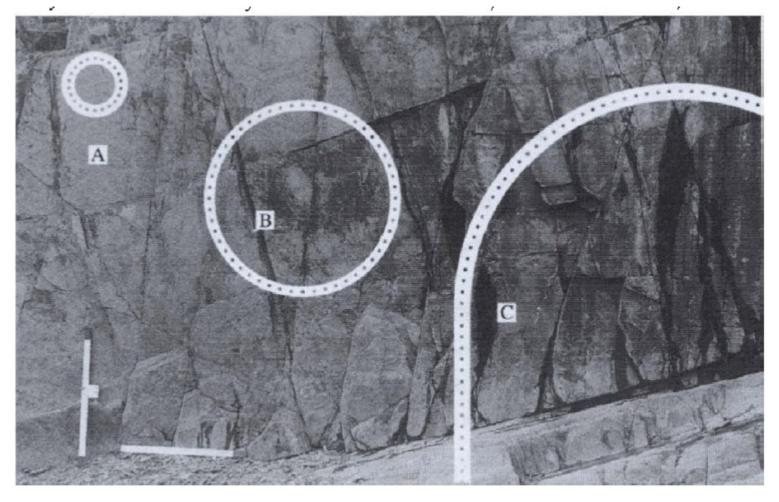
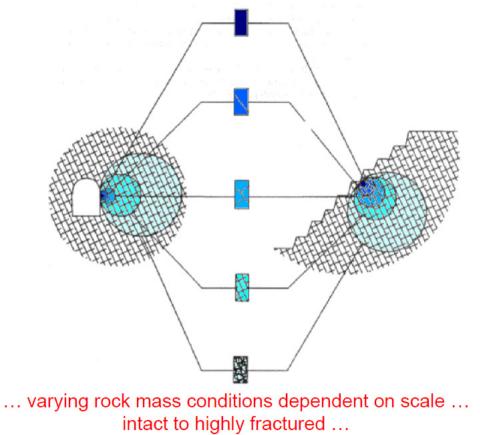


Figure 7.2 Discontinuities in rock and the engineered structure (A, borehole; B, tunnel; C, large excavation).

Discontinuity properties of rocks

Scale Effects in Rock Masses



The properties of discontinuities and rock mass classifications will be discussed in detail in following weeks (e.g. 5, 6, 8,9 and 11)

SEE YOUNEXT WEEK